



EARTHQUAKE

INTRODUCTION

- **Earthquakes** constitute one of the worst natural hazards which often turn into disaster causing widespread destruction and loss to human life.
- The effects of **earthquake** vary upon the magnitude and intensity. **Earthquakes** occur every now and then all round the world, except in some places where earthquakes occur rarely. The devastation of cities and towns is one of the effects of earthquake.



What is Earthquake?

An Earthquake is the result of a **sudden release of energy** in the earth's crust that creates seismic waves.

The seismic activity of an area refers to the **frequency, type and size of earthquakes experienced over a period of time**



DEFINITION

Earthquake = Vibration of the Earth produced by the rapid release of energy

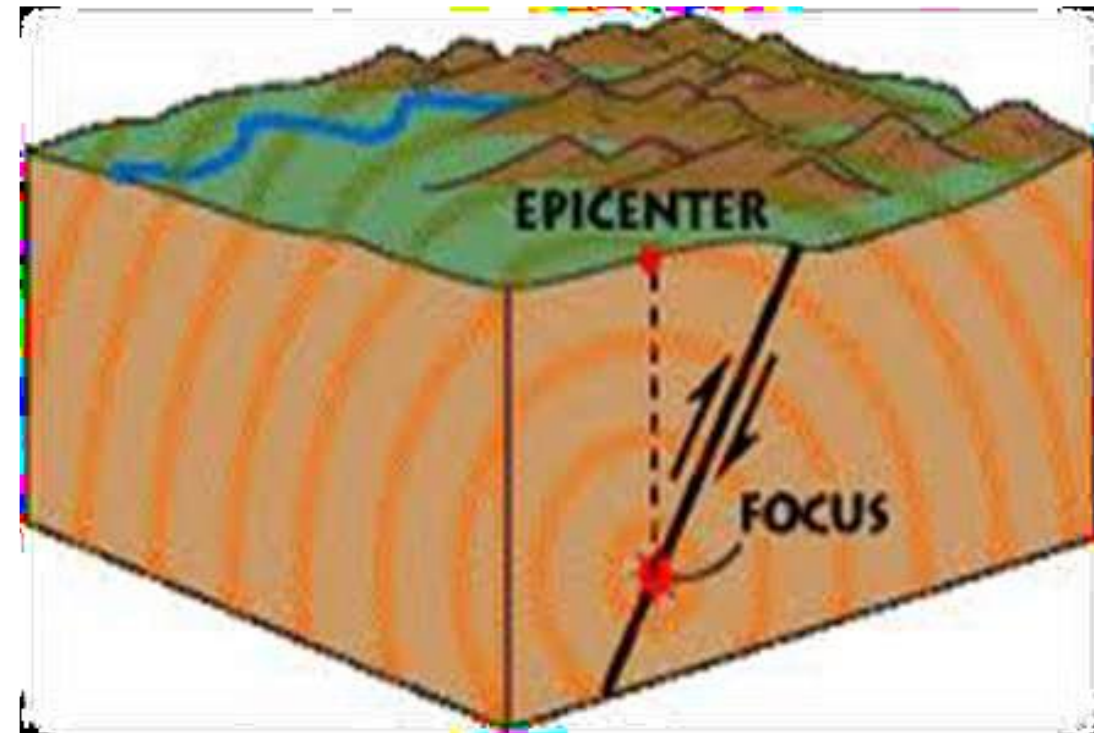
Seismic waves = Energy moving outward from the focus of an earthquake

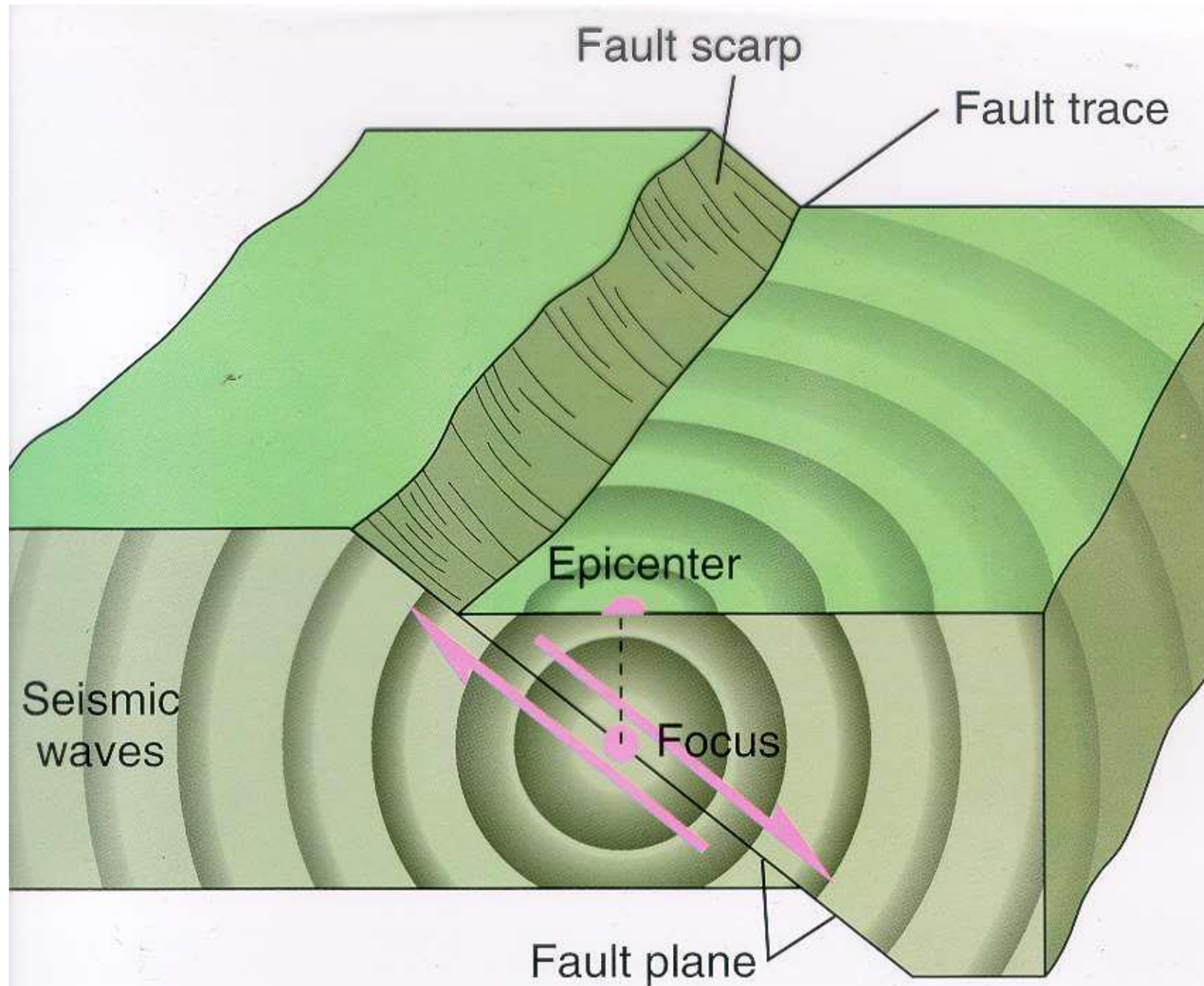
Focus(Hypocenter):

Focus is the point on the fault where rupture occurs and the location from which seismic waves are released.

Epicenter:

Epicenter is the point on the earth's surface that is directly above the focus ,the point where an earthquake or underground explosion originates.





Causes Of Earthquake

The primary cause of an earthquake is **faults** on the crust of the earth.

“A Fault is a break or fracture b/w two blocks of rocks in response to stress.”

- This movement may occur rapidly, in the form of an earthquake or may occur slowly, in the form of creep.
- Earth scientists use the angle of the fault with respect to the surface (known as the dip) and the direction of slip along the fault to classify faults.

Some major causes of earthquakes on basis of its causes are:

- Surface causes
- Volcanic causes
- Tectonic causes

Surface cause:

Great explosions, landslides, slips on steep coasts, dashing of sea waves , avalanches , railway trains, heavy trucks, some large engineering projects cause minor tremors. some of them are man made, other are natural.

Volcanic cause:

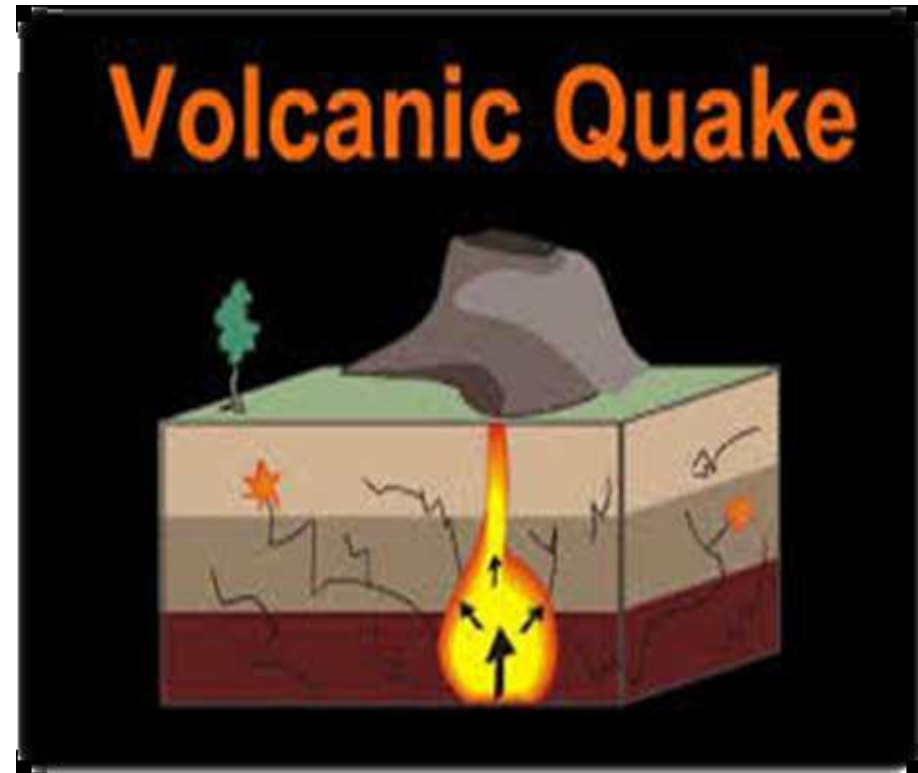
Volcanic eruptions produce earthquakes. Earthquakes may precede, accompany and frequently follow volcanic eruptions.

They are caused by sudden displacements of lava within or beneath the earth crust.

There are two general categories of earthquakes

that can occur at a volcano:

- volcano-tectonic earthquakes**
- long period earthquakes**



Tectonic cause:

Structural disturbances resulting in the parts of the lithosphere is the main cause of this type of earthquake.

Most of the disastrous earthquakes belong to this category and occur in areas of great faults and fractures.

Sudden yielding to strain produced on the rocks of accumulating stress causes displacements especially along old fault zones known as great transform faults.

Waves produced due to Earthquake

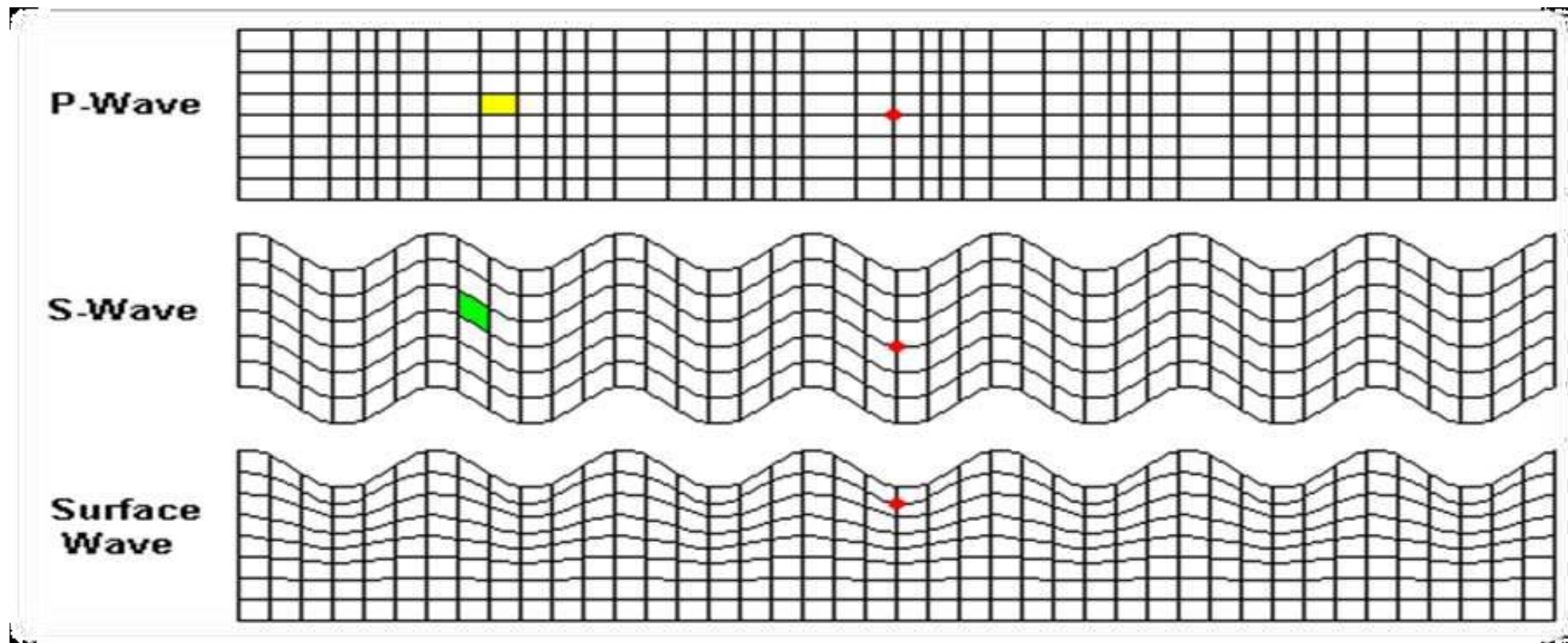
Seismic waves produced due to earthquake are basically divided into two major types:

- Body waves
- Surface waves

Body waves:

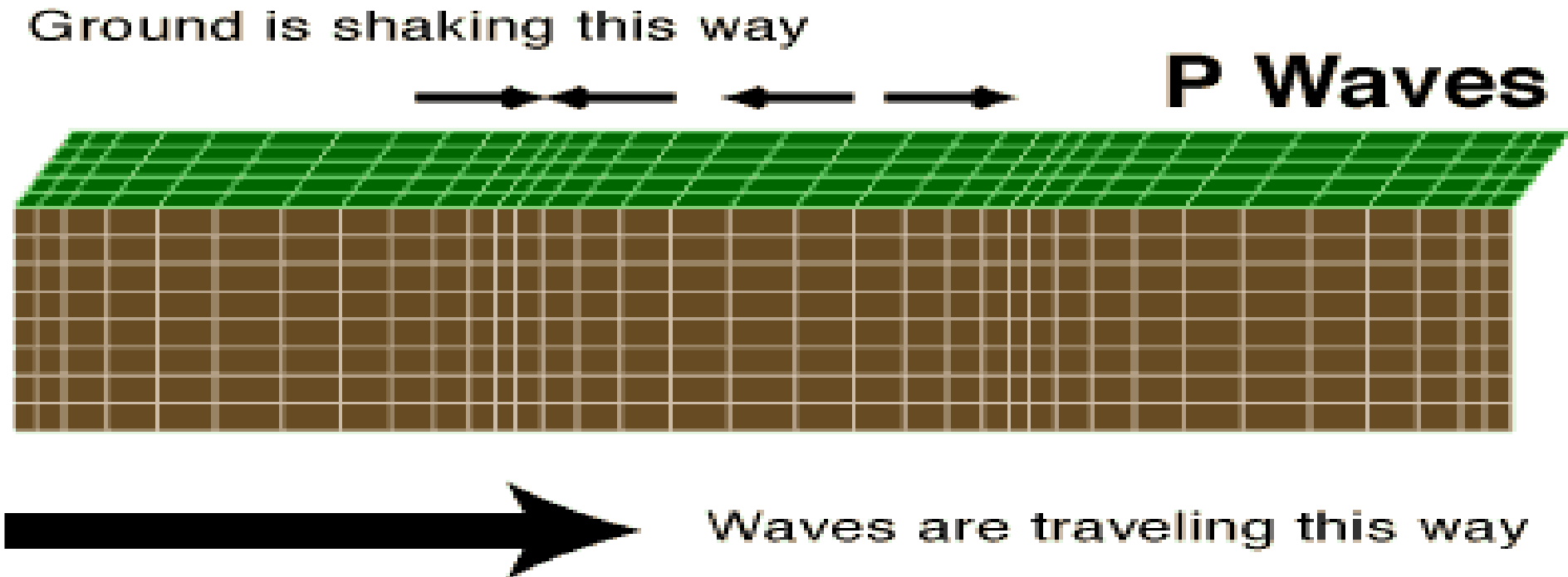
Body waves travels through the interior(body) of earth as they leave the focus. Body waves are further divided into following types:

- Primary (P) waves
- Secondary(S) waves



Primary Waves:

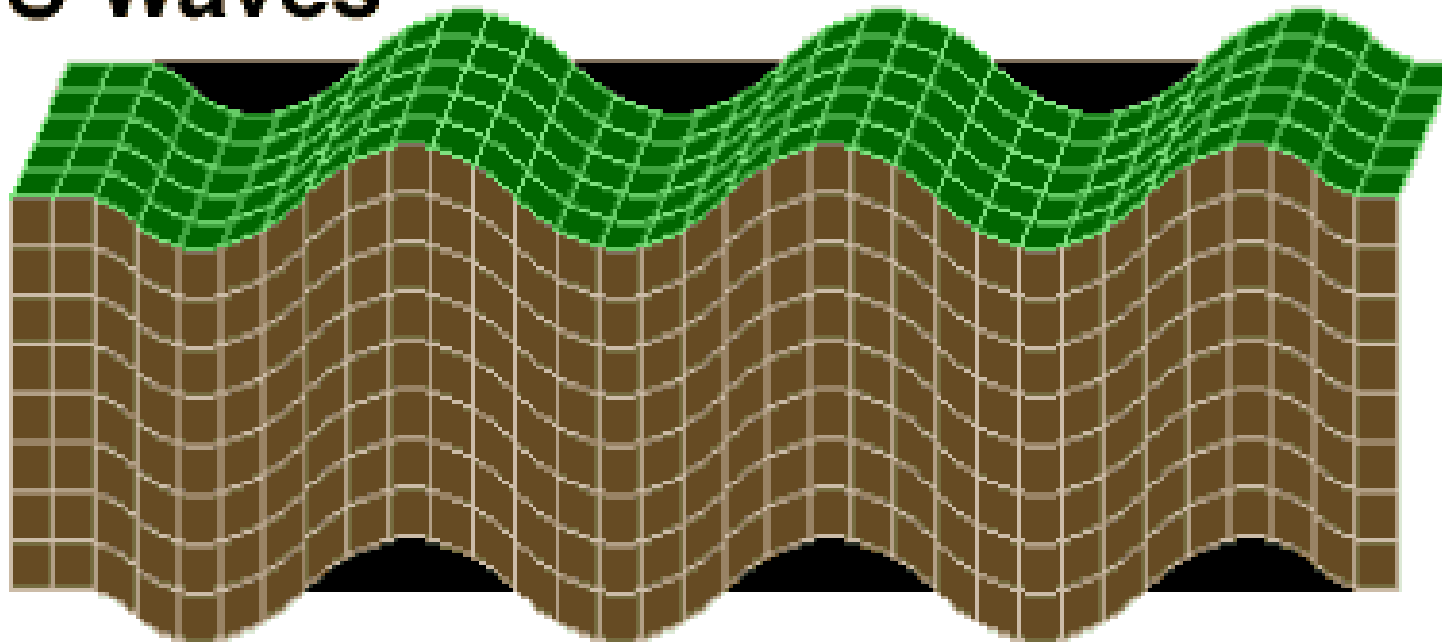
- called compressional, or push-pull waves.
- Propagate parallel to the direction in which the wave is moving.
- Move through solids, liquids



Secondary Waves (S);

- Called shear waves.
- Propagate the movement perpendicular to the direction in which the wave is moving.

S Waves



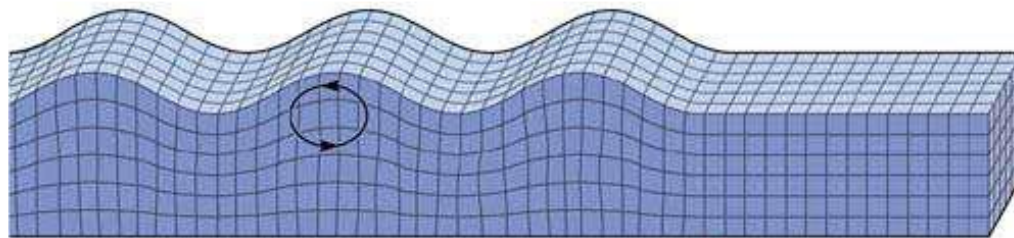
Primary Waves (P-waves)	Secondary Waves(S-wave)
High frequency	High frequency
Short Wavelength	Short Wavelength
Longitudinal waves	Transverse waves
Pass trough both solids and liquids	Can not move through liquids
Move forwards and backwards as it compressed and decompressed	Move in all direction from their source
P-wave is faster	S-wave is more slower than P-wave
First P-wave arrive	After P-wave,S-wave is arrive

Surface Wave:

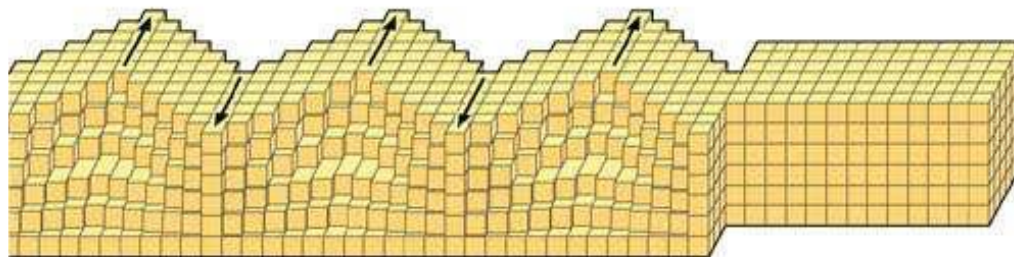
Surface waves travel parallel to the earth's surface and these waves are the slowest and most damaging.

Surface waves are divided into the following types:

- Love waves
- Rayleigh waves



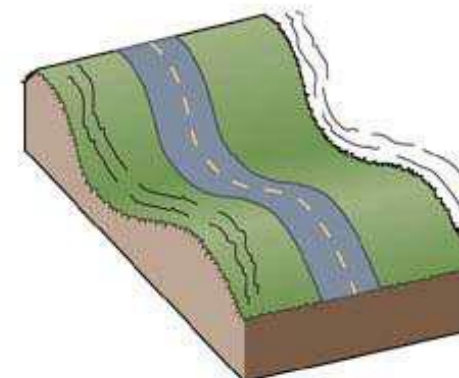
(a) Rayleigh wave



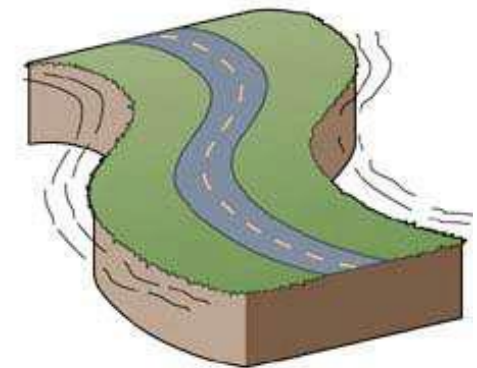
(b) Love wave



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Rayleigh wave



Love wave

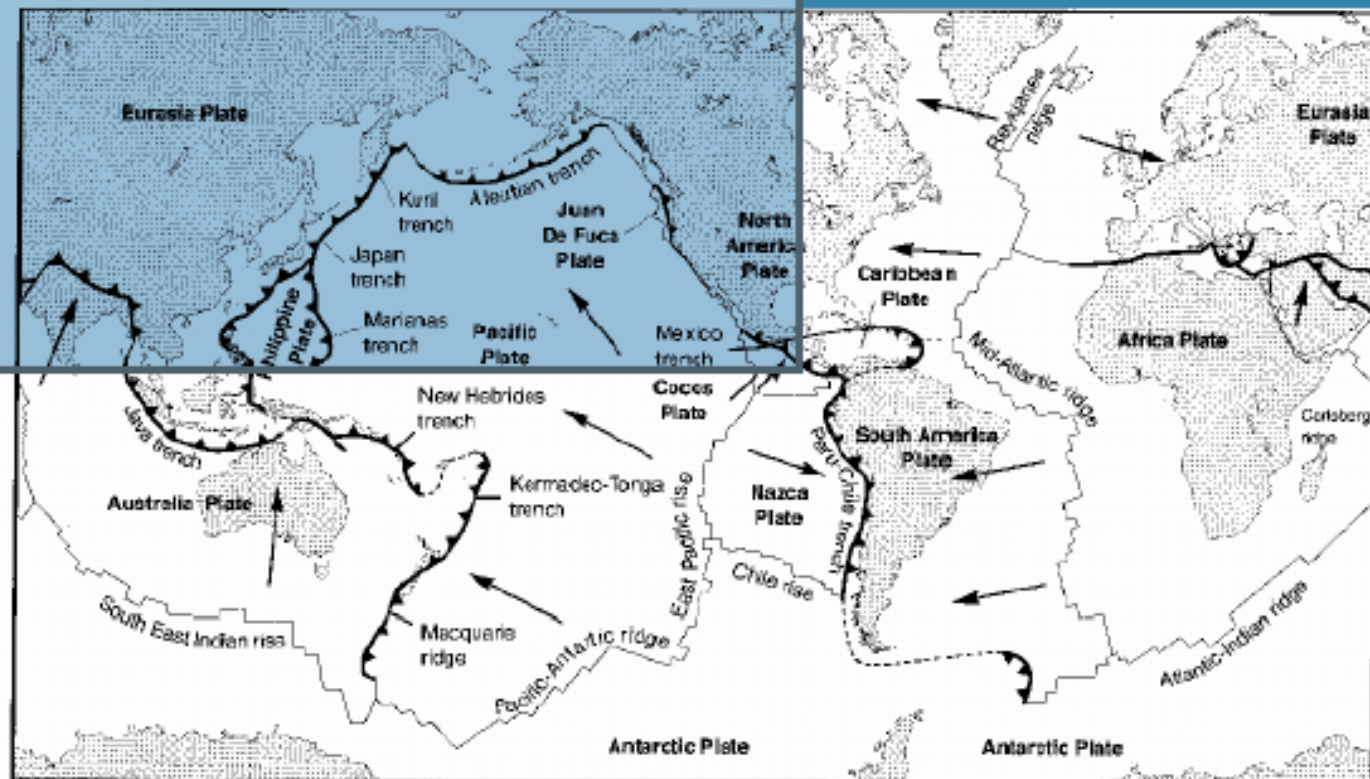
(c)

Love Waves	Rayleigh wave
Guided waves	Guided waves
Displacement is parallel to the free surface	Displacement is perpendicular to love-wave displacement
Love wave is faster	Rayleigh wave is slower
Causes horizontal shifting of the earth surface.	Ground move in circular motion.

Plate Tectonics

The earth's crust is divided into six continental-sized plates

- African
- American
- Antarctic
- Australia-I
- Eurasian
- Pacific



Key

Subduction zone	Uncertain plate boundary
Strike-slip (transform) faults	Ridge axis

Convection Currents in Mantle

- Near the bottom of the crust, horizontal component of convection currents impose shear stresses on bottom of crust, causing movement of plates on earth's surface.
- The movement causes the plates to move apart in some places and converge in others.

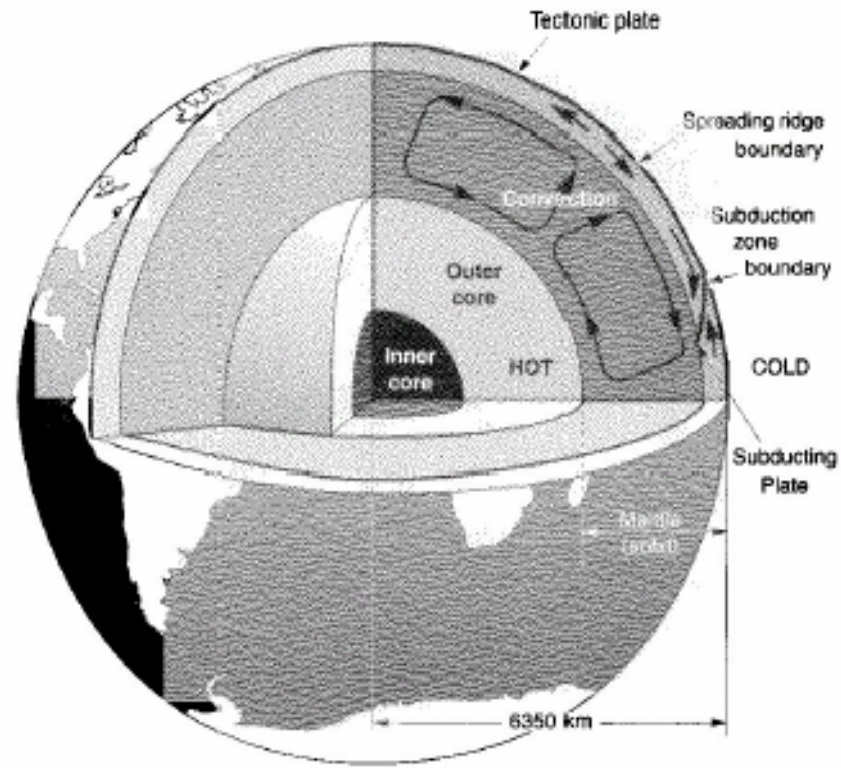


Plate Tectonics: The crust in motion

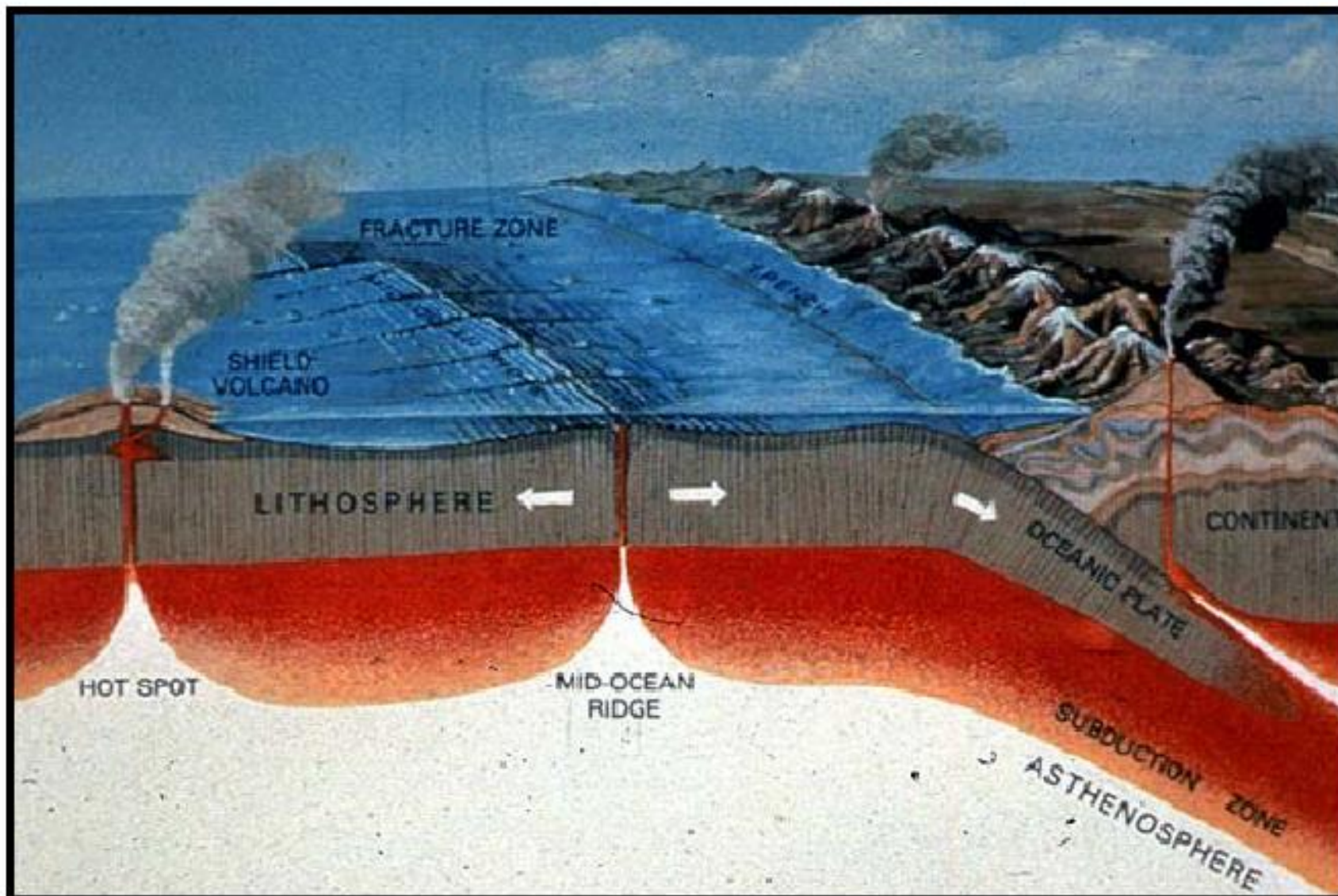
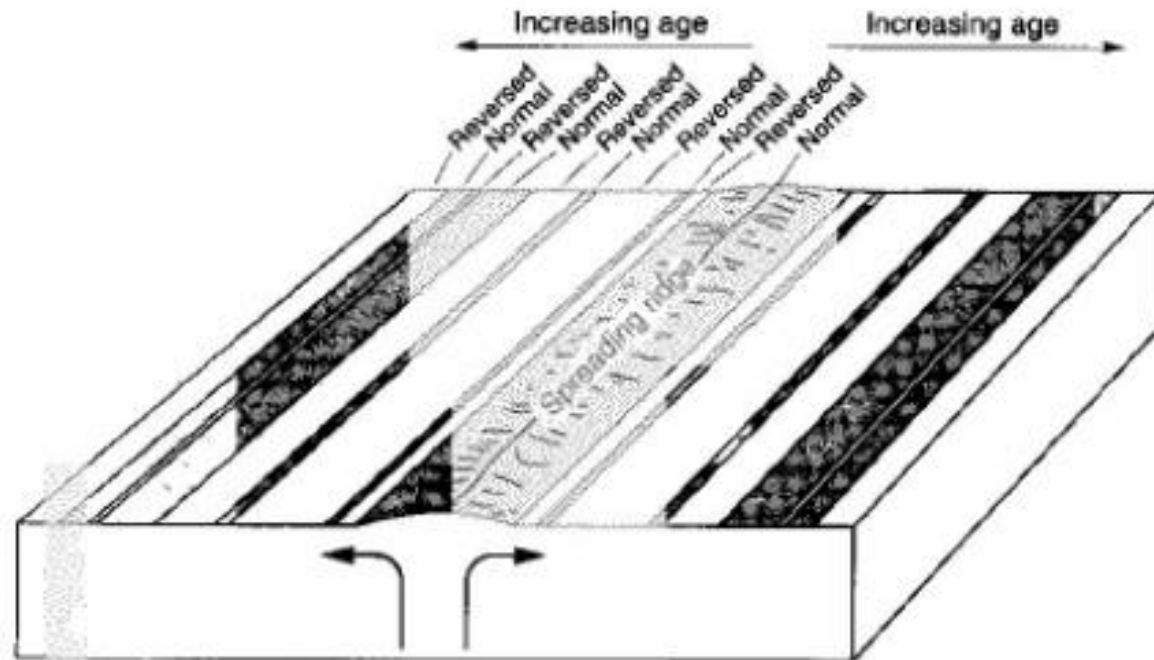


Photo courtesy of the USGS

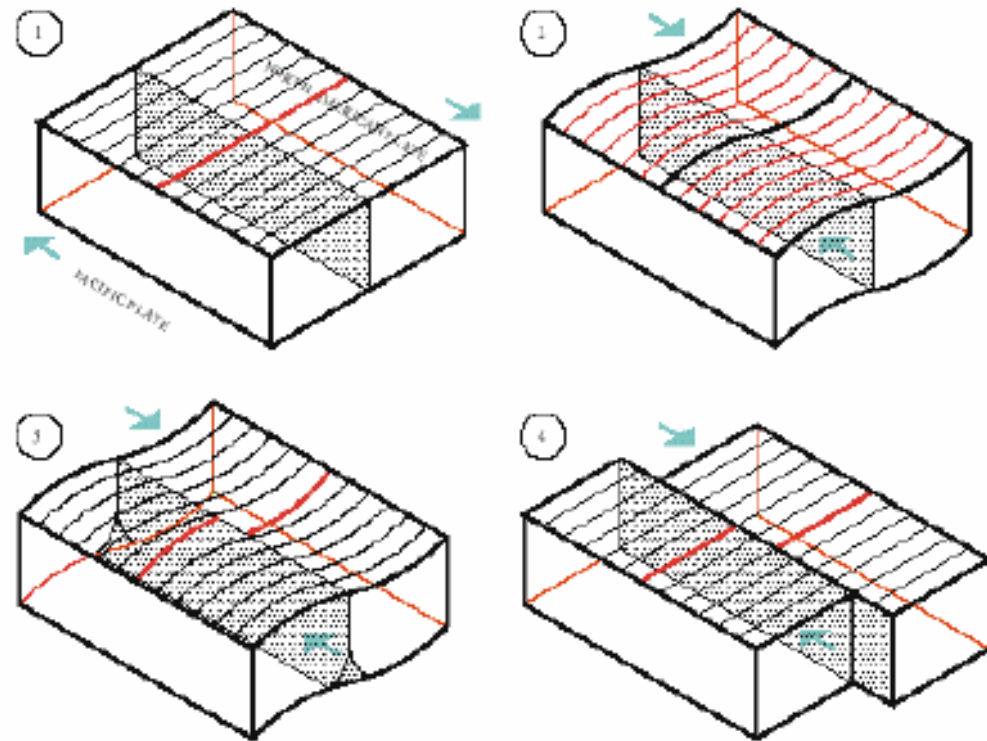
Spreading Ridge Boundaries

- Magma rises to surface and cools in gap formed by spreading plates.
- Magnetic anomalies are shown as stripes of normal and reversed magnetic polarity



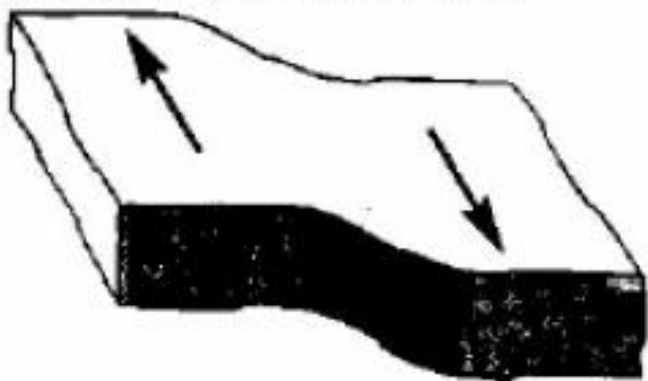
Elastic Rebound Theory of Earthquakes

Rocks are elastic, and mechanical energy can be stored in them just as it is stored in a compressed spring. When the two blocks for the opposite sides of the fault move by a small amount, the motion elastically strains rocks near the fault. When stress becomes larger than the frictional strength of the fault, the fictional bond fails at its weakest point. That point is the initial rupture, called the hypocenter, which may be near the surface or deep below it.



Elastic Rebound Theory of Earthquakes

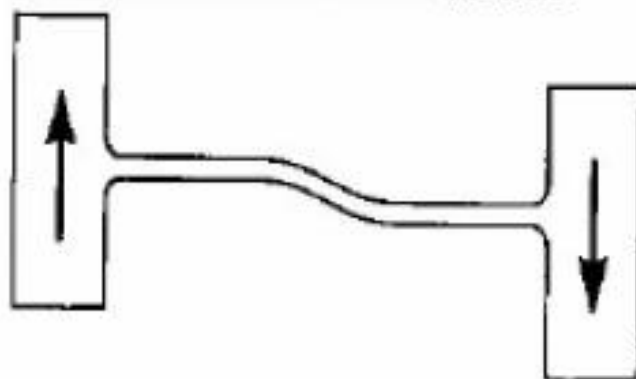
Deformation of ductile rock



Fracture of brittle rock

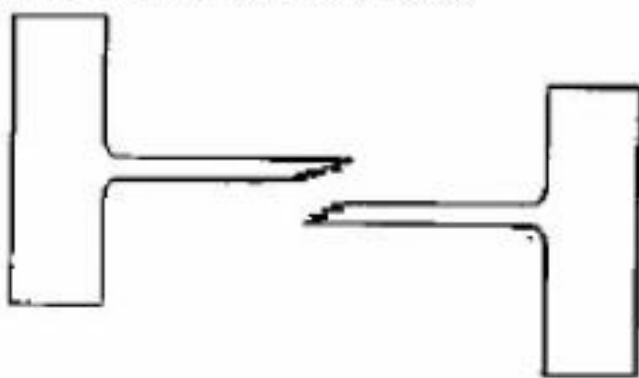


Deformation of a ductile stick



(a)

Fracture of a brittle stick



(b)



Magnitude and Intensity

Magnitude is directly related to energy release due to plate movement. It has a unique value for an earthquake.

It varies from **1** to **10**.

Intensity is related to *human feelings, behavior of secondary structures and structural behavior*. It has different values at different locations for a particular earthquake.

It varies from **I** to **XII**.

Magnitude and Intensity of Some Historical Earthquakes in Bangladesh

Event Name	M	I _{Dhaka}	Distance
1869 Cachar	7.5	V	250 km
1885 Bengal	7.0	VII	170 km
1897 Great Indian	8.7	VIII+	230 km
	8.0	Ambraseys, 2000	
1918 Srimangal	7.6	VI	150 km
1930 Dhubri	7.1	V+	250 km

Trifunac and Brady (1975): $\log(\text{PGA}) = 0.014 + 0.3 * I$

EMS VIII

- a) Many people find it difficult to stand, even outdoors.
- b) Furniture may be overturned. Waves may be seen on very soft ground.
- c) Many buildings of **class B (Unreinforced brick)** and a few of **class C (Unreinforced with RC floors/Frame with no antiseismic design)** suffer damage of **grade 3**.






Many buildings of **class A (Mud house)** and a few of **class B** suffer damage of **grade 4**; a few buildings of **class A** suffer damage of **grade 5**.

VULNERABILITY TABLE

Type of Structure		Vulnerability Class					
		A	B	C	D	E	F
MASONRY	rubble stone fieldstone	○					
	adobe (earth brick)	○—					
	simple stone	—○					
	massive stone	—○—					
	unreinforced brick / concrete blocks	—○—					
	unreinforced brick with RC floors	—○					
	reinforced brick (confined masonry)	—○—					
REINFORCED CONCRETE (RC)	RC without antiseismic design (ASD)	—○—					
	RC with minimum level of ASD	—○					
	RC with moderate level of ASD	—○					
	RC with high level of ASD				—○		
WOOD	wooden structures	—○—					
	+ Bamboo	—○—					

DAMAGE GRADE

Table 2: Classification of damage to masonry buildings

 <p style="text-align: center;">Grade 1</p>	<p>Grade 1: Negligible to slight damage (no structural damage) hair-line cracks in very few walls; fall of small pieces of plaster only. Fall of loose stones from upper parts of buildings in very few cases only.</p>
 <p style="text-align: center;">Grade 2</p>	<p>Grade 2: Moderate damage (slight structural damage, moderate non-structural damage) cracks in many walls; fall of fairly large pieces of plaster; parts of chimneys fall down.</p>
 <p style="text-align: center;">Grade 3</p>	<p>Grade 3: Substantial to heavy damage (moderate structural damage, heavy non-structural damage) large and extensive cracks in most walls; pantiles or slates slip off. Chimneys are broken at the roof line; failure of individual non-structural elements.</p>
 <p style="text-align: center;">Grade 4</p>	<p>Grade 4: Very heavy damage (heavy structural damage, very heavy non-structural damage). serious failure of walls; partial structural failure.</p>
 <p style="text-align: center;">Grade 5</p>	<p>Grade 5: Destruction (very heavy structural damage) total or near total collapse.</p>

Richter Magnitude

- In 1935, Charles Richter used a Wood-Anderson seismometer to define a *magnitude* scale for shallow, local (epicentral distances less than about 600 km (375 miles)) earthquakes in southern California.
- The Richter Magnitude, M , is calculated from the maximum amplitude, A , of the seismometer trace (Wood-Anderson Seismometer, $T_0 = 0.8\text{sec}$ and $\sigma = 0.80$) at a distance of 100 km from the epicenter.

$$M = \log A$$

however, a standard seismometer is not always at 100 km from the epicenter, in which

$$M = \log A - \log A_0$$

A_0 = maximum recorded amplitude for a particular earthquake selected at a site, generally $A_0 = 0.001$ mm for 100 km distance.